### Loki Modeling Update



#### Robert R. Howell University of Wyoming



### Two Component Models

- $T_{cool}$  (1-f) of surface
- $T_{hot}$  f of surface
- 3 variables, 3 data points 2.44, 3.28, 4.70µm but fits other wavelengths well too
- Using "image cube" (not tube) so some spatial mixing. Be suspicious of "edge" effects
- Similar to work in Davies 2003









Age Based on T<sub>cool</sub>

Wave velocity ~0.94 km inner ~1.38 km/day outer

Roughly ½ the Rathbun value of ~2 km/day

Related to lower flux during "non-periodic" I32 time?

T<sub>cool</sub> Black=260K White=300K

 $T = 538 K (t / days)^{-1/8}$ 

190 -



Temperature controls

- Why is temperature distribution this smooth?
  Differences of a degree or two over 10's of km
- Why are discrete temperatures present in different regions of the "two intensity" plots?



### Two Wavelength NIMS Plots



- Use <u>simultaneously</u> measured fluxes from I32 NIMS "Tube"
  - avoids spatial registration and averaging problems
- Plot 3.0007 μm flux vs. 4.6967 μm flux
- Model as single hot temperature (T) filling fraction of pixel (f)
  - Because wavelengths are relatively close, T and f are not "orthogonal"
  - Solid straight lines = constant T (from 300K to 700K)  $(F_v/F_x \Leftrightarrow T)$
  - Dashed curve lines = constant f (from 1.0 to 0.0001)
- Colors "arbitrary", used to identify points from different regions of plots for later use
  - However, even without colors, points clearly fall into distinct regions of 2I plot



## Large version of 21 plot



### Spatial Pattern of Clusters



- Red: Hot western edge (with hottest at very edge)
- Aqua and yellow: Inner and outer margins
- Orange: Location near "island"
- Green: Bulk of patera with smooth spatial gradient from "resurfacing wave"



#### Additional structure within the "Green Band"





- Red: Hot western edge (with hottest at very edge)
- Aqua and yellow: Inner and outer margins
- Orange: Location near "island"
- Green: Bulk of patera with smooth spatial gradient from "resurfacing wave"



# Possible buffering by volatiles

- 200 days old ⇒277K ⇒366 W/m<sup>2</sup> radiative flux from surface
   presumably balanced by conductive flux from below
- To produce this from S vapor deposition (latent heat)  $\Rightarrow 0.036$  m/day
  - Some reasonable fraction of this needed to buffer temperatures
- Depth consistent with geometric contraints
- Is it consistent with constraints from
  - Spectra
  - Required pressure
  - Sources of S



### Vapor transport models

- Adapted from Ingersoll (and others) SO<sub>2</sub> models
- Expect transport on horizontal scale of  $L = (2 \pi)^{1/2} (k T)/(mg) = 15 \text{ km}$
- Consider 300 K vs. 260 K surface separated by 15 km
  - S vapor pressure at 300K is few nanobars
  - Produces transport and deposition too low by factor of ~100
- At 350K S vapor pressure is high enough to produce sufficient transport
  - Because of exponential dependence of P on T
  - No 350K present in "low T" component of surface
  - 350K is present only 0.3 m below the surface
  - Even higher temperatures present in the high T "crack" fraction
- Just getting started on transport models involving small hot component / vertical transport



### Spectral-Morphological Constraints

• Only high resolution images are from Voyager





#### C10 is best Galileo full wavelength set





#### C10 spectra normalized to "island"





- Analysis of morphology / temperature structure in NIMS data (2I plots)
- Better modeling of volatile transport
- Spectral analysis
  - Proper normalization (true reflectance) for Galileo spectra
  - Comparison with standards and other Io regions
  - Incorporation of higher resolution Voyager images

