that the troughs are a truly global phenomenon. No trough was seen near longitude 270°W, but that region is geologically complex, with numerous cross-cutting dark lineaments. The north-south offset of the antipodal centers of symmetry, maintained in the New Horizons images, hints at true polar wander of Europa’s ice shell (21–23).

From Earth, the solar phase angle $g$ for the Jupiter system is $<12°$, limiting Earth-based measurement of directional scattering by jovian satellites. Only spacecraft can access higher phase angles. LORRI observed the Galilean satellites at a range of phase angles (table S1), filling gaps in Europa’s solar phase curve between 32° and 103° and between 109° and 143°. The photometric data, corrected for longitudinal variations and normalized to Voyager data (24), are shown in Fig. 4. Europa’s brightness at $g = 70°$ is more than 40% of its fully illuminated brightness, underscoring the unique texture of its surface produced by active resurfacing. The comparable number for Earth’s Moon is only 20%.

Our observations improve measurement of Europa’s phase integral $q$, which describes the directional scattering properties of light reflected from its surface. The new $q$ value is $1.01 \pm 0.04$, compared with $1.1 \pm 0.1$ from previous Voyager data (24). Compared with other actively resurfaced icy satellites, Europa’s $q$ is marginally higher than that of Enceladus ($0.89 \pm 0.10$) (25) but lower than that of Triton $(1.14 \pm 0.03)$ (26). Using a geometric albedo of $0.67 \pm 0.03$ for LORRI wavelengths ($\lambda \geq 24$), we find a new Bond albedo of $0.68 \pm 0.05$, compared with a previous value of $0.6 \pm 0.1$ (24), meaning that Europa absorbs less sunlight than previously thought.

References and Notes

28. We thank the entire New Horizons mission team and our colleagues on the New Horizons science team. New Horizons is funded by NASA whose financial support we gratefully acknowledge. We also thank C. A. Hibbitts, J. B. Dalton III, G. B. Hansen, and K. Stephan for valuable scientific discussions as well as providing examples of NIMS data.

Supporting Online Material

www.sciencemag.org/cgi/content/full/318/5848/234/DC1
Table S1
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REPORT

Io’s Atmospheric Response to Eclipse: UV Aurorae Observations


The New Horizons (NH) spacecraft observed Io’s aurora in eclipse on four occasions during spring 2007. NH Alice ultraviolet spectroscopy and concurrent Hubble Space Telescope ultraviolet imaging in eclipse investigate the relative contribution of volatiles to Io’s atmosphere and its interaction with Jupiter’s magnetosphere. Auroral brightness and morphology variations after eclipse ingress and egress reveal changes in the relative contribution of sublimation and volcanic sources to the atmosphere. Brightnesses viewed at different geometries are best explained by a dramatic difference between the dayside and nightside atmospheric density. Far-ultraviolet auroral morphology reveals the influence of plumes on Io’s electrodynamic interaction with Jupiter’s magnetosphere. Comparisons to detailed simulations of Io’s aurora indicate that volatiles supply 1 to 3% of the dayside atmosphere.

Io is a volcanically active moon of Jupiter, and its volcanism is the ultimate source of material for Io’s sulfur-dioxide atmosphere. The interaction between Io’s atmosphere and the Io plasma torus produces displays of auroral emissions on Io, supplies plasma to Jupiter’s magnetosphere, and physically links Io to Jupiter (J). The relative importance of the volcanoes as a direct, immediate source of the atmosphere, versus sublimation of frosts deposited around these volcanoes, has remained uncertain since the atmosphere’s discovery in 1979 (2, 3). Io’s average dayside surface temperature rapidly drops after eclipse ingress or at night (likely from ~120 K to ~90 K; (4, 5)), which is sufficient to diminish the sublimation component of the atmosphere across most of the surface and possibly results in an atmosphere mostly supplied directly from volcanoes.
New Horizons at Jupiter

stream relative to the magnetospheric plasma flow). These large-scale features are known to change brightness and location with Jupiter’s changing magnetic field orientation at Io and are diagnostic of the local flow of the Io plasma torus past the satellite and into its atmosphere (10, 11, 20–23). Dramatic visible auroral glows have been seen from numerous volcanic plumes, including the large Tyssdrur plume that was active at this time (17).

The question of how much of the dayside atmosphere comes from SO2 frost sublimation versus volcanoes remains difficult to resolve. Previous theoretical work demonstrated that auroral brightness variations with time after Io enters eclipse ingress provide a means to investigate the relative contributions of volcanic and sublimation sources to Io’s dayside atmosphere (14). A relative contribution from volcanoes of 1 to 10% was suggested based on only a few data points with inadequate time coverage.

We observed the time series of Io’s FUV emissions in shadow using Alice (Fig. 1) and the HST/SBC (Fig. 2). The last Alice exposure in IEclipse05 (red points in Fig. 3, A to C) is ~40% brighter than it is in earlier measurements (Table 2) and likely represents predicted postmidaurora brightening (14). In IEclipse01, the view is of the dayside, similar to that for HST (Fig. 2). This view includes Io’s wake emissions. The high initial brightnesses and the decrease in brightness after ingress in IEclipse01 (Fig. 3) may be because the amount of sublimation is particularly sensitive to changes after eclipse ingress on the dayside atmosphere. There was little change in volcanic activity observed during the encounter period (17); thus, any changes in Io’s neutral atmospheric density must be attributed to other sources. The plasma torus density is thought to be relatively stable on the time scale of days. The trend of brighter aurora with Io’s location in denser regions of the plasma torus (21) would exacerbate the difference in measured OI 135.6-nm brightnesses between IEclipse01 and IEclipse04 (see Fig. S1B). Viewing geometry of the asymmetric atmosphere regions and the large-scale auroral features viewed (e.g., wake viewed in IEclipse01 but not IEclipse04) could explain the differences, but the FUV equatorial spots are consistently brighter than IEclipse04 (Io radii), respectively, centered on Io. The most prominent FUV emission lines that contribute to this image are the same neutral OI 130.4 nm, OI 135.6 nm, and SIl 147.9 nm lines observed with Alice (Fig. 1B). However, the first SBC exposure occurred 12 min after unural ingress, which is after the time expected for the most dramatic variations. Alice observed the key periods for IEclipse01 and IEclipse04, but only two points at ~90 min after ingress for this IEclipse03 event. Cassini visible aurora imaging (11) showed similarly moderate variations in brightness over a period starting ~20 min after ingress.

A comparison between aurora simulations and the auroral brightness time series shown in Fig. 3 enables a higher fidelity assessment of Io’s atmospheric sources. Volcanic column densities of 2 × 10^{14} cm^{-2}, 4 × 10^{14} cm^{-2}, and 8 × 10^{14} cm^{-2} out of 1.5 × 10^{16} cm^{-2} were used in Fig. 3 (1%, 3%, and 5% cases, respectively). The data are best described by a volcanic contribution of 1 to 3% to a primarily sublimation-supplied dayside atmosphere.

Table 1. NH Io eclipse Alice observation summary. Supporting observations with the HST/SBC are also indicated. IEclipse02 was dropped from the plan and not performed.

<table>
<thead>
<tr>
<th>Visit name</th>
<th>Date</th>
<th>Umbral ingress</th>
<th>Umbral egress</th>
<th>Instrument used</th>
<th>Number of exposures</th>
</tr>
</thead>
<tbody>
<tr>
<td>IEclipse03</td>
<td>2/27/2007</td>
<td>14:21</td>
<td>16:27</td>
<td>Alice</td>
<td>2</td>
</tr>
<tr>
<td>IEclipse04</td>
<td>3/1/2007</td>
<td>8:50</td>
<td>10:56</td>
<td>Alice</td>
<td>20</td>
</tr>
</tbody>
</table>

The HST/SBC data (Fig. 3D) show that Io’s FUV brightness decreased between the first and last exposures by roughly factors of 1.5 and 1.3 for square regions of widths 0.55 R_Io and 4 R_Io (Io radii), respectively, centered on Io. The most prominent FUV emission lines that contribute to this image are the same neutral OI 130.4 nm, OI 135.6 nm, and SiI 147.9 nm lines observed with Alice (Fig. 1B). However, the first SBC exposure occurred 12 min after unural ingress, which is after the time expected for the most dramatic variations. Alice observed the key periods for IEclipse01 and IEclipse04, but only two points at ~90 min after ingress for this IEclipse03 event. Cassini visible aurora imaging (11) showed similarly moderate variations in brightness over a period starting ~20 min after ingress.

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Other methods for characterizing the atmospheric sources face the predicament of collocated volcanic plumes and SO2 frost deposits, potentially confusing spatial associations of higher density near volcanoes with the respective sources. Recent reports using these other methods are in general agreement with the results of 1 to 3% reported here (25, 26), but SO2 maps derived from Lyman-α absorption images show more gas near the equatorial terminators than expected (27, 23, 25) and remain difficult to reconcile with these levels of sublimation.

The latitudinal extent of the sub-Jupiter equatorial spot (Fig. 2, C to F) decreased by ~25% from 17 to 37 min after unural ingress. Space Telescope Imaging Spectrograph (STIS) Io eclipse observations similarly hinted at FUV emissions more localized to the magnetic equator (13). In shadow, preferential closure of Io’s electrodynamic current system was predicted to occur through local plume atmospheres (29, 30). A combination of local dominance of the plume atmosphere over the sublimation source could explain this variation in the latitudinal extent of the sub-Jupiter equatorial spot with time in eclipse.

An emission feature between 200 km and 550 km above the limb and located near the bright, newly discovered “East Girru” volcano is seen in
**Fig. 2.** HST/SBC FUV images of Io on 27 February 2007, concurrent with NH IEclipse03 observations. (**A**) Average of all eight 5-min HST exposures. (**B**) Same as (**A**), with labeled features. The sub-Earth longitude is 344°. The sub-jovian (0°) and orbital trailing (270°) longitudes are indicated with dashed and dotted meridians, respectively, with 15° gridding for the graticule. The view of Io from Earth is 33° from the concurrent NH view. Contours indicate 0.5 kilorayleigh (kR) increments in brightness levels. The aurora morphology is qualitatively consistent with previous FUV observations. The upstream-side auroral feature is nominally expected where the jovian magnetic field is tangent to the limb (shown with arrow). However, its location is ~15° North and likely influenced by the East Girru plume ([17]). The wake auroral feature is observed to extend >1 R João, downstream relative to the plasma flow. The 1.4 ± 0.1 times brighter southern polar limb glow measured between 0.75 and 1.25 R João is consistent with Io being located “above” the plasma torus centrifugal equator ([23]). The nearby north, south, and west Masubi volcanoes known to be active at this time ([17]) may also enhance the southern limb glow brightness on the left (fig. S2). The F125LP filter includes FUV emissions in the 125-nm to 200-nm band-pass, which are primarily OI 130.4 nm, OI 135.6 nm, Si 138.9 nm, Si 147.9 nm, Si 166.7 nm, Si 181.2 nm, and Si 190.0 lines for atomic gas, with smaller contributions from ionized Si 125.6 nm, Si 142.9 nm, and Si 172.9 nm ([20]).

**Fig. 3.** Time series of Io’s auroral emissions in Eclipse. (**A**) NH Alice IEclipse01, IEclipse03, IEclipse04, and IEclipse05 brightness measurements of OI 135.6 nm emissions are shown with time after umbral eclipse ingress. (**B**) Same as (**A**), but normalized to values in eclipse from 50 to 125 min after ingress and compared with aurora simulations ([24]) for three levels of volcanic contribution. (**C**) Same as (**A**) and (**B**), but normalized to pre-ingress, sunlit values. The last measurement in IEclipse05 supports the predicted post-eclipse brightening. IEclipse01 views the dayside, whereas IEclipse04 views mostly the nightside and is dimmer. These diurnal (phase angle) variations indicate that the atmosphere in shadow (both on the nightside and in eclipse) is supplied primarily by volcanoes (see additional plots in fig. S1). IEclipse01, obtained at roughly twice the distance from NH as IEclipse04, has more statistical noise. (**D**) HST/ACS/SBC brightness measurements of regions within the limb and extending a few R João with time in eclipse, concurrent with the IEclipse03 event. The SBC brightnesses decreased by a factor of 1.3 during the period between 20 min and 50 min after ingress.
Table 2. Alice-measured emission line brightness averages and SDs in sunlight and eclipse.

<table>
<thead>
<tr>
<th>Emission line</th>
<th>Type</th>
<th>IEclipse01</th>
<th>IEclipse04</th>
<th>IEclipse05</th>
</tr>
</thead>
<tbody>
<tr>
<td>OI 130.4 nm</td>
<td>Sunlight</td>
<td>706 ± 134</td>
<td>445 ± 27</td>
<td>347 ± 142</td>
</tr>
<tr>
<td></td>
<td>Eclipse</td>
<td>597 ± 43</td>
<td>394 ± 14</td>
<td>254 ± 71</td>
</tr>
<tr>
<td></td>
<td>Ratio S/E</td>
<td>1.18 ± 0.24</td>
<td>1.12 ± 0.08</td>
<td>1.37 ± 0.68</td>
</tr>
<tr>
<td>OI 135.6 nm</td>
<td>Sunlight</td>
<td>882 ± 177</td>
<td>577 ± 35</td>
<td>480 ± 188</td>
</tr>
<tr>
<td></td>
<td>Eclipse</td>
<td>797 ± 57</td>
<td>536 ± 18</td>
<td>361 ± 94</td>
</tr>
<tr>
<td></td>
<td>Ratio S/E</td>
<td>1.11 ± 0.24</td>
<td>1.08 ± 0.08</td>
<td>1.26 ± 0.58</td>
</tr>
<tr>
<td>Sl 147.9 nm</td>
<td>Sunlight</td>
<td>1167 ± 271</td>
<td>986 ± 54</td>
<td>596 ± 288</td>
</tr>
<tr>
<td></td>
<td>Eclipse</td>
<td>1205 ± 87</td>
<td>874 ± 28</td>
<td>429 ± 144</td>
</tr>
<tr>
<td></td>
<td>Ratio S/E</td>
<td>0.97 ± 0.24</td>
<td>1.13 ± 0.07</td>
<td>1.39 ± 0.82</td>
</tr>
</tbody>
</table>

NH Long Range Reconnaissance Imager (LORRI) images (17). This same feature appears in the HST/SBC image in Fig. 2B, obtained when East Girru was shifted just behind the limb. The auroral feature near East Girru in both LORRI and HST/SBC images is ~15° northward of Jupiter’s field line tangent point at the limb, which suggests that ionospheric currents are diverted northward from this nominal position toward a region of higher gas density near the plume. Similar deviations of the anti-Jovian FUV emissions from nominal tangent points observed with STIS are likely caused by the prevalence and distribution of plumes there (21). Volcanic plume aurorae were not identified in previous lower-quality STIS FUV images, which caused an apparent discrepancy with visible images of plume aurorae. The East Girru plume FUV auroral feature in Fig. 2 resolves this discrepancy and reveals the influence of plumes on Io’s electrodynamic interaction. The upstream-side emission feature is more apparent when limb brightened at viewing geometries like those reported in Fig. 2. This feature was predicted by aurora image simulations (22) and is diagnostic of the divergence of the plasma flow upstream of Io, a primary trait of Io’s interaction with the plasma torus.

References and Notes
18. The angular size of Io varies with spacecraft distance but is smaller than the Alice slit width for these data. The spectral resolution varies between 0.3 nm and ~0.9 nm for emissions known to be located near the satellite disk (22); see, e.g., Fig. 2.

REPORT

Io Volcanism Seen by New Horizons: A Major Eruption of the Tvashtar Volcano

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Jupiter’s moon Io is known to host active volcanoes. In February and March 2007, the New Horizons spacecraft obtained a global snapshot of Io’s volcanism. A 350-kilometer-high volcanic plume was seen to emanate from the Tvashtar volcano (62°N, 122°W), and its motion was observed. The plume’s morphology and dynamics support nonballistic models of large Io plumes and also suggest that most visible plume particles condensed within the plume rather than being ejected from the source. In images taken in eclipse, nonthermal visible-wavelength emission was seen from individual volcanoes near Io’s sub-Jupiter and anti-Jupiter points. Near-infrared emission from the brightest volcanoes indicates minimum magma temperatures in the 1150–1335-kelvin range, consistent with basaltic composition.

The New Horizons (NH) Jupiter flyby provided the first close-up observations of the tidally driven volcanism of Jupiter’s moon Io since the last Galileo orbiter observations of Io in late 2001 (1). The closest approach to Io occurred at 21:57 UT on 28 February 2007 at a range of 2.24 mrad per pixel Multicolor Visible Imaging Camera (MVIC), although MVIC coverage included those at Zal and Kuradalagon and a large new plume, 150 km high, at northern Lema Regio, which has produced a large albedo change. Three of these plumes, north Lema and north and south Masubi, are associated with recent large lava flows, supporting the idea that Prometheus-type plumes result from mobilization of surface volatiles by active lava flows. All active plumes that were on...